

Adaptive Multiscale Streamline Simulation and Inversion for High-Resolution Geomodels

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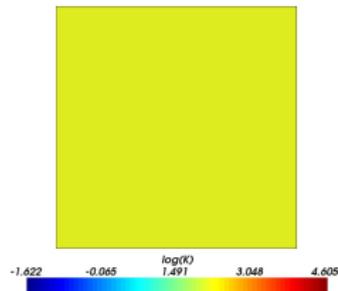
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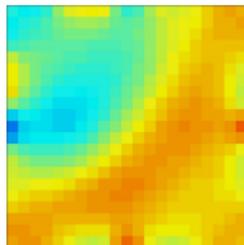
Introduction: History matching

History matching is the procedure of modifying the reservoir description to match measured reservoir responses.

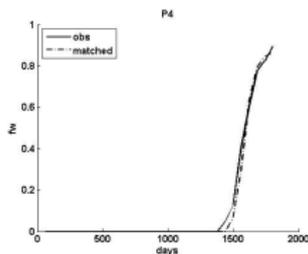
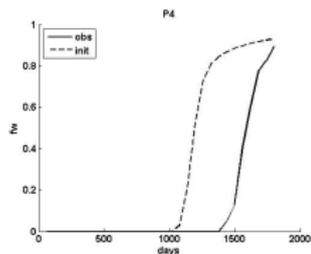
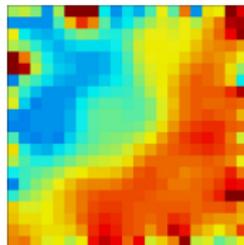
Initial:



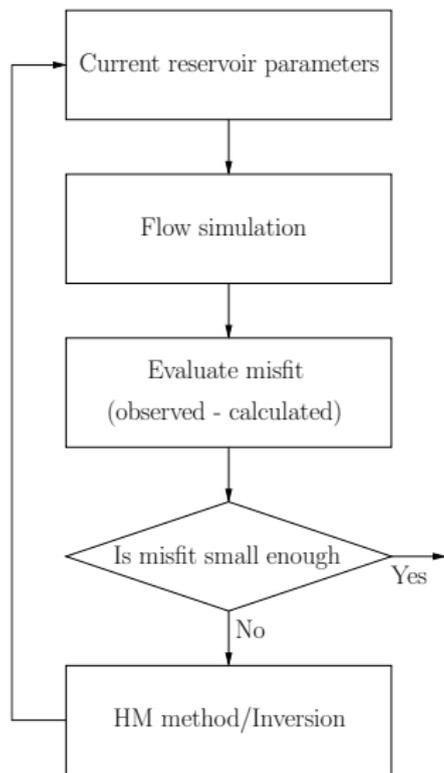
Matched:



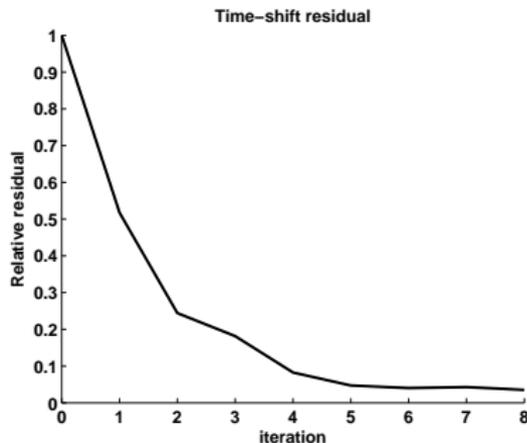
Reference:



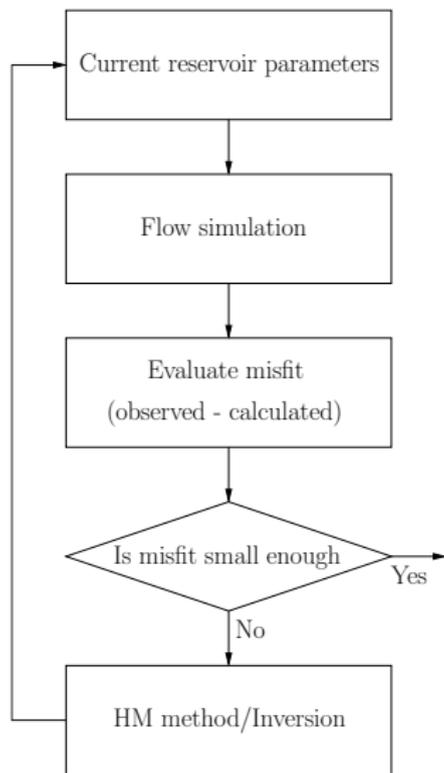
Introduction: History-matching loop



$$E = \sum (d^{\text{obs}} - d^{\text{cal}})^2, \quad d^{\text{cal}} = g(\mathbf{m})$$



Challenges in history-matching loop



Problems:

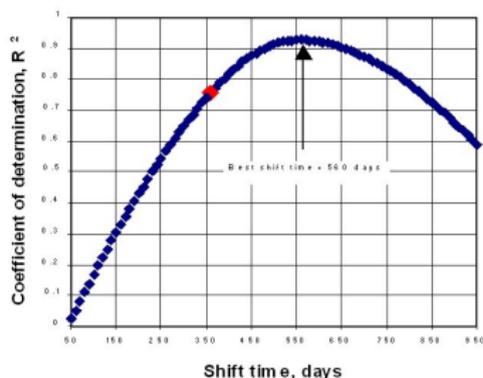
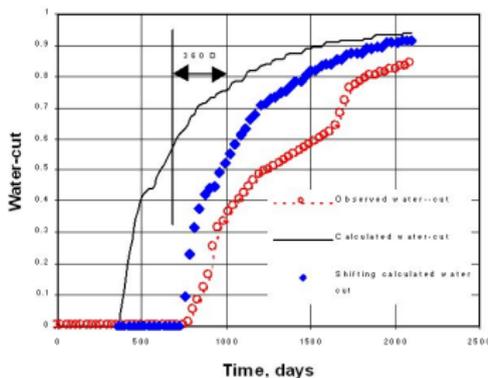
- highly under-determined problem → non-uniqueness
- errors in model, data, and methods
- nonlinear forward model
- *non-convex misfit functions*
- *forward simulations are computationally demanding*

Challenge I: Non-convex misfit function

Inversion method: Generalized Travel-Time (GTT) inversion with analytic sensitivities [Vasco et al. (1999), He et al. (2002)]

The generalized travel time is defined as the 'optimal' time-shift that maximizes

$$R^2(\Delta t) = 1 - \frac{\sum [y^{\text{obs}}(t_i + \Delta t) - y^{\text{cal}}(t_i)]^2}{\sum [y^{\text{obs}}(t_i) - \bar{y}^{\text{obs}}(t_i)]^2}.$$



Basic underlying principles for the history–matching algorithm

- Minimize travel-time misfit for water–cut by iterative least-square minimization algorithm.
- Preserve geologic realism by keeping changes to prior geologic model minimal (if possible).
- Only allow smooth large-scale changes. Production data have low resolution and cannot be used to infer small-scale variations.

Minimization of functional:

$\Delta\tilde{\mathbf{t}}$: Travel–time shift

\mathbf{S} : Sensitivity matrix

\mathbf{m} : Reservoir
parameters

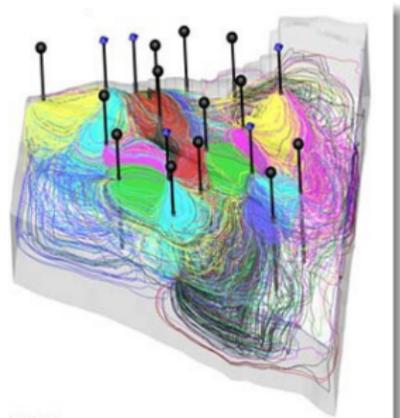
$$\|\Delta\tilde{\mathbf{t}} - \mathbf{S}\delta\mathbf{R}\| + \overbrace{\beta_1\|\delta\mathbf{R}\| + \beta_2\|\mathbf{L}\delta\mathbf{R}\|}^{\text{Regularization}}$$

norm *smoothing*

\mathbf{S} computed analytically along streamlines from a single flow simulation

Features of streamlines

- Very well suited for modeling large heterogeneous multiwell systems dominated by convection
- Generally fast flow simulation
- Delineate flow pattern (injector-producer pairs)
- Enables analytic sensitivities



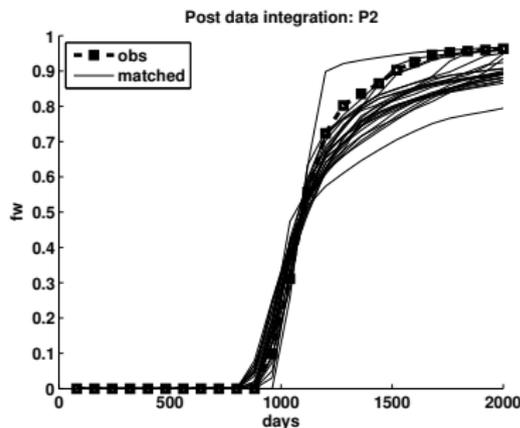
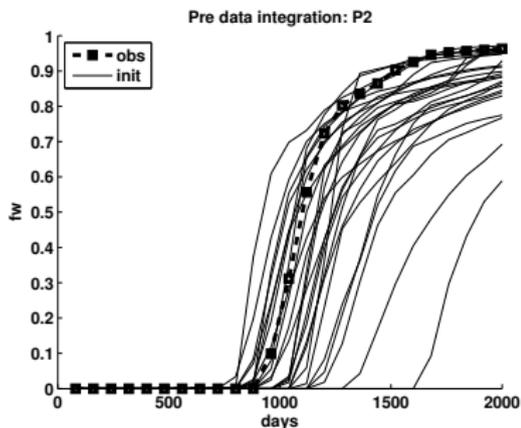
Source: www.techplot.com

Classes of streamline-based history-matching methods

- Assisted history matching
- (Generalized) travel-time inversion methods
- Streamline-effective property methods
- Miscellaneous

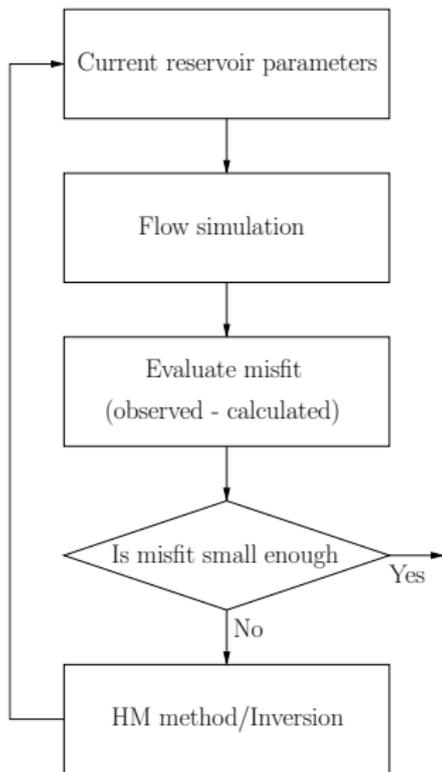
Example: Uncertainty quantification

Simple two-phase model (end-point mobility $M = 0.5$) on a 2D horizontal reservoir, 25×25 cells with lognormal permeability. Result after eight iterations:



Statistical analysis of mean and standard deviation

Challenge II: Long runtime for forward simulations



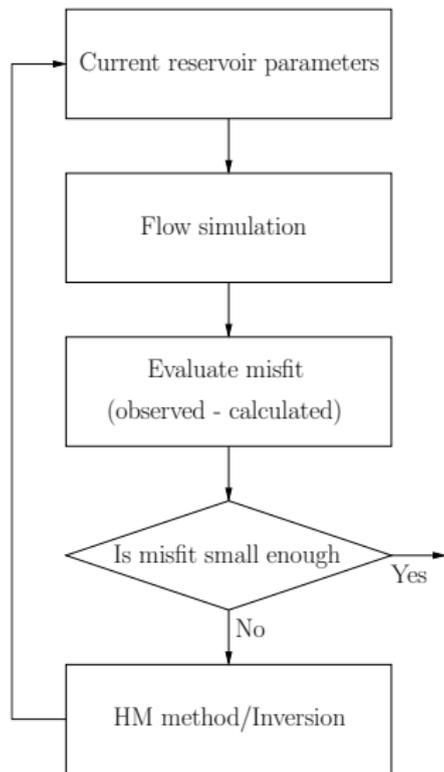
Streamline simulation much faster than conventional FD-methods.

Still, room for improvement.

Observations:

- pressure solver most expensive part of flow simulation
- parameters change very little from one flow simulation to the next

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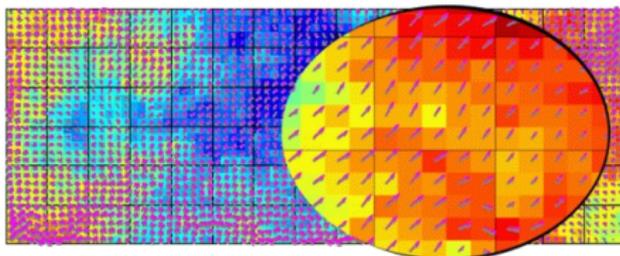
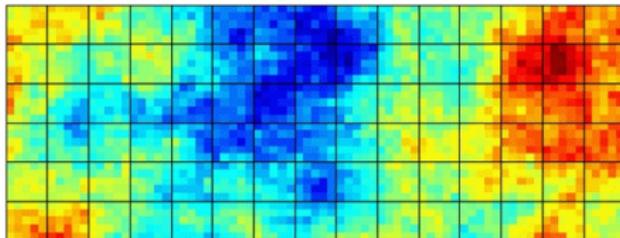
- pressure solver most expensive part of flow simulation
- parameters change very little from one flow simulation to the next

Idea: should reuse computations in areas with minor changes
→ multiscale methods

Multiscale pressure solver

Upscaling and downscaling in one step. Runtime like coarse-scale solver, resolution like fine-scale solver.

Fine grid: 75×30 . Coarse grid: 15×6

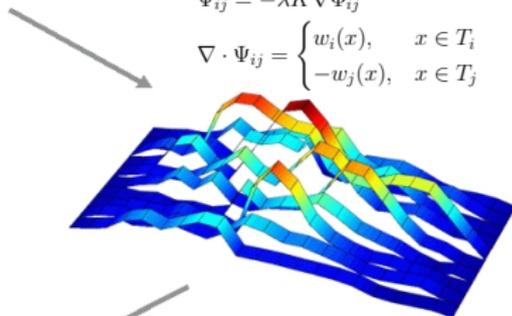


Coarse grid: pressure and fluxes. Fine grid: fluxes

Basis functions for each pair of coarse blocks $T_i \cup T_j$:

$$\Psi_{ij} = -\lambda K \nabla \Phi_{ij}$$

$$\nabla \cdot \Psi_{ij} = \begin{cases} w_i(x), & x \in T_i \\ -w_j(x), & x \in T_j \end{cases}$$

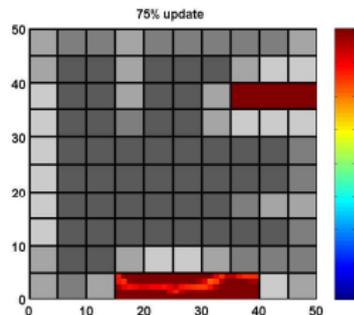
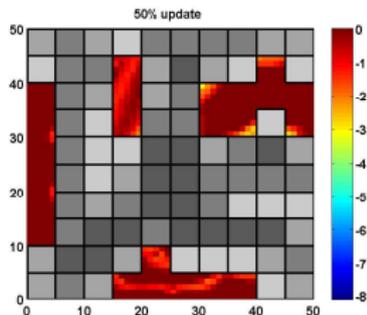
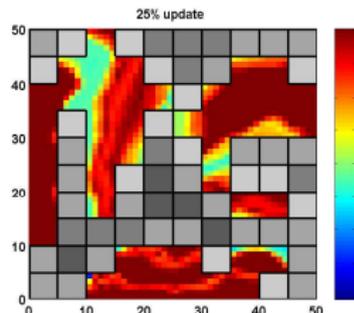
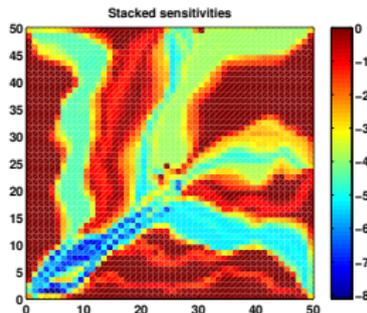
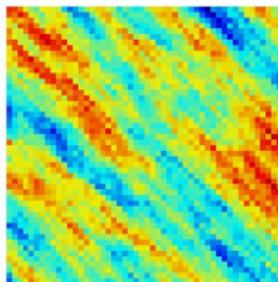


Global linear system with 249 unknowns:

$$\nabla \cdot v = q, \quad v = -\lambda K \nabla p$$

Further computational savings

Can also reuse basis functions from previous forward simulation.
General idea: use sensitivities to steer updating

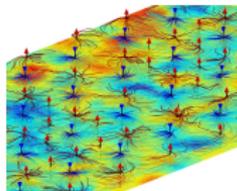


History matching on geological models

Generalized travel-time inversion on million-cell model

1 million cells, 32 injectors, and 69 producers
2475 days \approx 7 years of water-cut data

Analytical sensitivities along streamlines + travel-time inversion (quasi-linearization of misfit functional)



| Solver | Misfit | | | CPU-time (wall clock) | | |
|------------|--------|-------|---------------------------|-----------------------|--------|---------|
| | T | A | $\overline{\Delta \ln k}$ | Total | Pres. | Transp. |
| Initial | 100.0 | 100.0 | 0.821 | — | — | — |
| TPFA | 8.9 | 53.5 | 0.806 | 64 min | 33 min | 28 min |
| Multiscale | 11.2 | 47.3 | 0.812 | 43 min | 7 min | 32 min |
| Multiscale | 10.4 | 45.4 | 0.828 | 17 min | 7 min | 6 min |

Time-shift misfit: $\|\Delta \mathbf{t}\|_2$

Amplitude misfit: $[\sum_k \sum_j (d_w^{\text{obs}} - d_w^{\text{cal}})^2]^{1/2}$

Permeability discrepancy: $\sum_{i=1}^N |\ln k_i^{\text{ref}} - \ln k_i^{\text{match}}|/N$

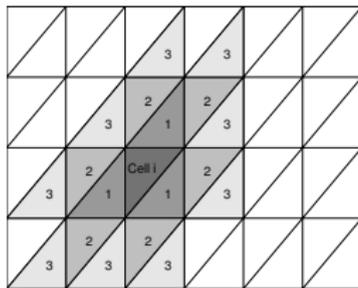
Extension to unstructured grids

Primary concern: how to define smoothing stencil

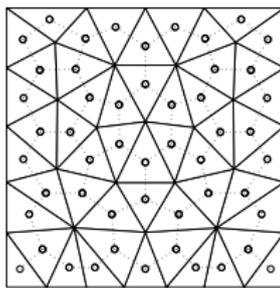
Generalized Laplacian stencil

$$L_i \mathbf{m} = -w_{ii}m_i + \sum_{j \in \mathcal{N}(i)} w_{ji}m_j,$$
$$w_{ji} = w_{\text{norm}} \cdot \rho(\zeta(i, j); R), \quad w_{ii} = \sum_{j \in \mathcal{N}(i)} w_{ji}.$$

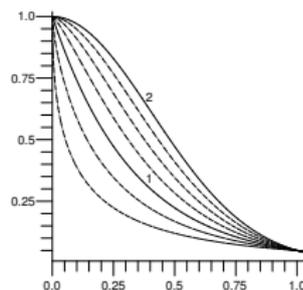
Neighborhood: $\mathcal{N}(i)$



Distance: $\zeta(i, j)$



Correlation: $\rho(\zeta; R, \dots)$



Properties of smoothing: independent of grid density, reduce to Laplacian on Cartesian grids, decay with distance $\zeta(i, j)$, be zero outside finite range, be bounded as $\zeta \rightarrow 0$.

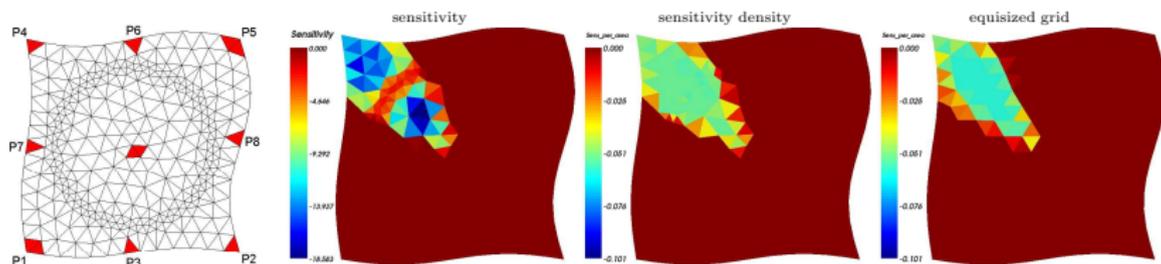
Extension to unstructured grids

Second concern: sensitivities

Sensitivities are spatially additive.

Small/large cells \longrightarrow small/large sensitivities

Thus, grid effects are to be expected



Rescaling of sensitivities:

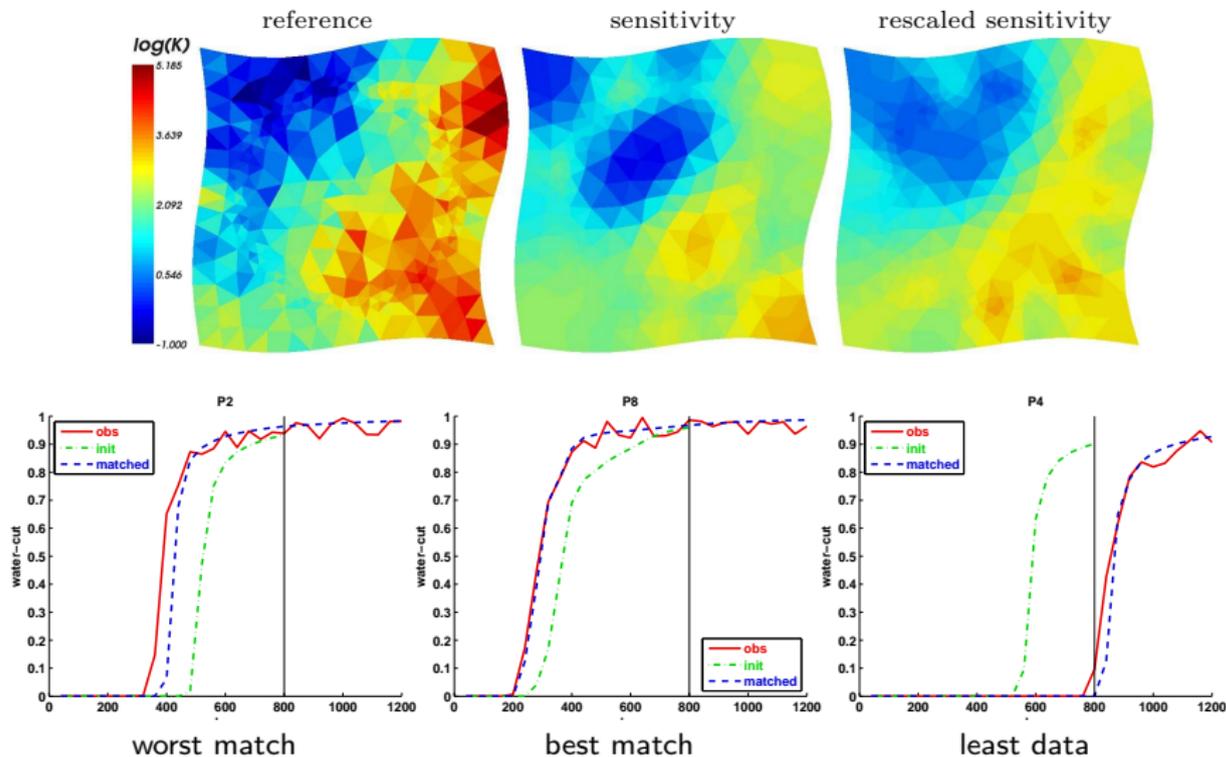
Permeability modification ΔK_i scales with sensitivity G_i .

Splitting cell in two $\implies \Delta K \rightarrow \sim \frac{1}{2} \Delta K$ in each subcell

Therefore: scale sensitivity by relative volume, $\tilde{G}_i = G_i(\bar{V}/V_i)$

Example:

1200 days observed (5% noise added), 800 matched, 400 predicted



Example:

Grid effects for models with thin layers

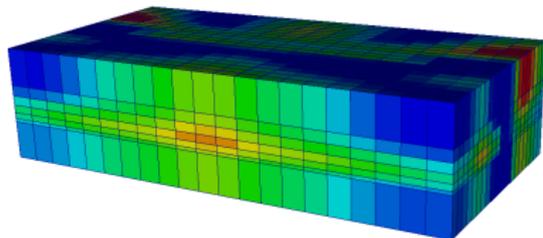
Synthetic test case:

$21 \times 21 \times 7$ tensor-product grid with layers of varying thickness.

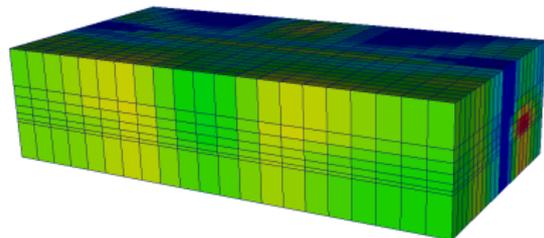
| Layer | 1 | 2 | 3 | 4 | 5 | 6 | 7 |
|-----------|-----|-------|-------|-------|-------|-------|-----|
| Thickness | 1.0 | 0.089 | 0.164 | 0.212 | 0.251 | 0.285 | 1.0 |

Initial model: two different constant values

True model: homogeneous



K from original G_{ij}



K obtained from scaled G_{ij}

Example:

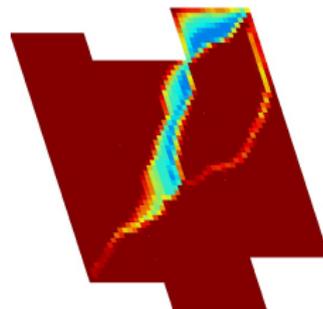
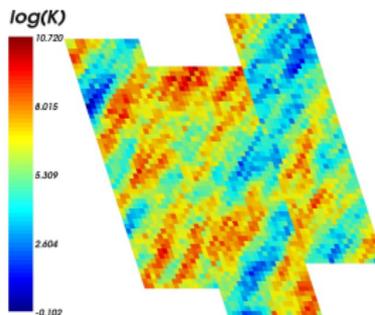
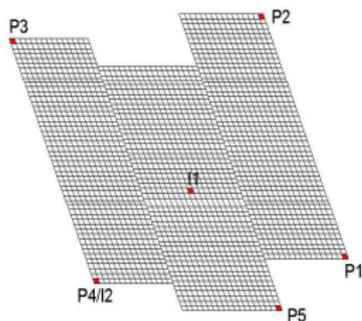
Corner-point grid with two non-sealing strike-slip faults

Infill drilling case:

Production data from 3000 days, 2500 used in history match.

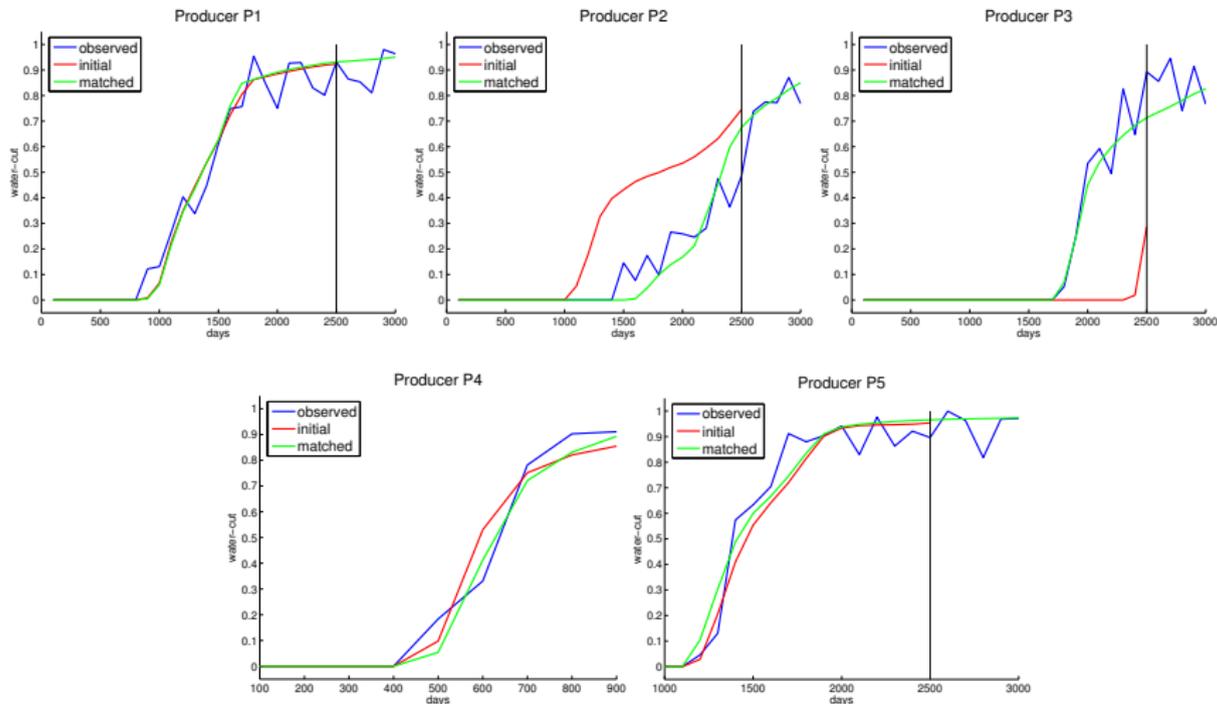
At 900 days:

- new producer P5 drilled
- producer P4 converted to injector



Example:

Corner-point grid with two non-sealing strike-slip faults, cont'd



Direct continuation of previous work:

- Unstructured grids (done for inversion algorithm)
- Corner-point grids (testing remains to be done on real models)
- Other types of data / more general flow

Other possible directions using streamlines:

- Closed-loop reservoir management
- History-matching seismic data
- Use of sensitivities for other optimization workflows
- ...